Astro 210 Lecture 18 March 2, 2011

Announcements

HW5 due at start of class Friday
 Oops! – typos found in Problem 2(d)
 which should refer to the situation in 2(c)
 erratum and corrected question posted

 Night Observing this week – *Dress warmly!* report forms, info online

Last time:

- atmospheres & properties of gasses

Last time: gasses under a microscope

gasses: collections of particles, e.g., atoms, molecules

www: gas law simulation

- in space: "elbow room" empty space between particles
- in time: constant random motion collisions exchange energy, momentum
- *individual* particle velocities random, changing but *distribution* (=histogram=bar graph) of particle speeds is set by temperature

 \rightarrow average speed $v_{\rm rms} = \sqrt{3kT/m}$

Microscopic View of a Piston

Now consider a large number of gas particles

- in a sealed volume
- with a *piston* of area A

from microscopic viewpoint: piston constantly bombarded by gas particles if let free—would be pushed away

to resist bombardment, must *push* on piston = exert *force* F define **pressure**

$$P = \frac{\text{force on piston}}{\text{area of piston}} = \frac{F}{A} \tag{1}$$

 $^{\omega}$ Q: how to intensify bombardment = pressure on piston? Hint-more than one way to do this

Pressure

collisions with walls \rightarrow momentum transfer \rightarrow force \rightarrow pressure www: piston simulation

ideal gas

pressure P, volume V, total number N of particles and absolute (Kelvin) temperature T all related by *ideal gas equation of state*:

PV = NkT

(2)

- N counts individual particles, typically very large! alternatively: can count in units of moles of particles i.e., in units of $N_{Avo} = 6 \times 10^{23}$ then # moles is $n_{moles} = N/N_{Avo}$ and $PV = n_{moles}RT$, where $R = N_{Avo}k = 8.3$ Joules mole⁻¹ K⁻¹
- since density $\rho = M/V = \mu N/V$, with μ gas particle mass can rewrite ideal gas law: $P = \rho kT/\mu$

Planetary Atmospheres

Terrestrial Atmospheres:

- atmospheres are tiny (or zero!) fraction of planet mass
- no light gasses (H₂, He), only heavier N₂, O₂, CO₂ (if anything!)

Jovian Atmospheres:

- a significant fraction of planet mass
- mostly H_2 , He; some heavier species

Q: why the difference? what factors important?

competition: gravity versus thermal motion

gravity \rightarrow keep particles thermal motion \rightarrow run away

• gravity
$$\rightarrow$$
 escape speed
 $v_{\rm esc} = \sqrt{2GM/R} = 2.4 \times 10^3$ m/s for Moon

• thermal motion \rightarrow avg thermal speed $v_{\rm rms} = 2.6 \times 10^3$ m/s for H on Moon $v_{\rm rms} > v_{\rm esc}$: many atoms can escape H lost from Moon (check!)

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really: not all particles have same speed so always some escape real question: timescale to keep atm for age of SS need $v_{esc} \ge 6v_{rms}$ (rule of thumb)

species	vesc $/v$ rms	lifetime
H on moon	0.9	hours
H on earth	2.1	days
sodium on moon	4	$\sim 10^3$ yrs
O ₂ on earth	12	$ ightarrow\infty$: no escape!

www: Sodium lunar ''atmosphere'' solar UV photon on Moon ''soil'' \rightarrow thin vapor of sodium \neg ejected but replenished



The Earth

astro-trivia: Earth logo/icon is \oplus recall: $R_{\oplus} = 6.4 \times 10^6$ m = 6400 km get mass from $g = GM_{\oplus}/R_{\oplus}^2$ (need G!) $\rightarrow M_{\oplus} = gR_{\oplus}^2/G = 6.0 \times 10^{24}$ kg

Average density:

 $ho_{\rm av,\oplus} = 3/4\pi M_\oplus/R_\oplus^3 = 5,500 \, {\rm kg/m^3}$ between rocks and iron \rightarrow some of both

Earth's Interior

crust: 16–40 km \rightarrow *thin!* mantle: ~ 3000 km outer core: ~ 2,200 km inner core: ~ 1,200 km

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crust: solid. ocean basins—basalt: O, Si, Al, Mg
continental plates–granite: O, Si, Al, Na, K
mantle: "plastic". Fe, Mg, Si, O
outer core: liquid. Fe, Ni
inner core: solid. Fe, Ni
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heaviest elements lowest \rightarrow settling ("differentiation")

Q: how do we know?

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How do we know?
"Refraction" of Earthquakes
Demo: slinky
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Earthquakes "emit" waves \rightarrow use seismographic info to learn earth structure

www: diagram

Plate Tectonics

- www: plot of earthquake sites
- www: plot of volcano sites
- www: plate locations

crust not a single rigid solid but collection of "plates" motions in mantle (convection) cause plate motion www: convection currents

www: satellite laser ranging
www: VLBI: radio telescopes used to detect motion
www: drift animation

plates move, and interact: collide, slide, buckle
⇒ "plate tectonics"
leads to observed geological features
www: Mountain
www: volcano (Kilauea, HI)

www: San Andreas

ex: San Andreas fault in CA: sliding plates \rightleftharpoons

Earth is evolving!

iClicker Poll: Earth's Orbit

Earth orbit around Sun: ellipse, eccentricity $e_{\oplus} = 0.017$

Below: one ellipse with $e = e_{\oplus}$, one circle e = 0Which of these has the eccentricity of Earth's ellipse?



Survey Says

Foci labeled



Seasons and Eccentricity

Due to elliptical orbit, Earth-Sun distance changes $r_{ap} = (1 + e)a_{\oplus} = 1.017 \text{ AU}$ $r_{peri} = (1 - e)a_{\oplus} = 0.983 \text{ AU}$...but not by much! 3.2% swing!

if this were the whole story: temperature swing by $\approx 2\% \times 300~\text{K} = 6^\circ~\text{C} = 11^\circ~\text{F!}$

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also: if yearly temperature variations were due to eccentricity...*Q: how should the seasons compare in the N and S hemispheres?Q: what is actually observed?Q: what does this imply about season origin?*

Origin of the Seasons

seasons are **exactly opposite** in N and S hemispheres i.e., right now is spring in S. America, will soon be summer

so: season origin must distinguish hemispheres

recall: Earth's spin axis tilted w.r.t. orbit by 23.5°

- when one hemisphere tilted closest, the other farthest: this is summer/winter
- when tilt \perp Earth-Sun radius: fall/spring

Note: Earth closest to Sun in northern winter! \rightarrow eccentricity has tiny effect on temperature